Titanomagnetite, a new potential geochronometer for in situ U-Pb dating

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Supplementary Information

Table S-1 The modified ID-TIMS U-Pb age results for zircons HGF-01 from the Hongge intrusion³¹.

Sample	²⁰⁷ Pb/ ²³⁵ U		²⁰⁶ Pb/ ²³⁸ U		
Number	Ratio	2σ	Ratio	2o	— rho
1	0.2906	0.0068	0.04102	0.0008	0.751
2	0.2914	0.0055	0.04099	0.0007	0.773
3	0.291	0.0066	0.04107	0.0008	0.746
4	0.2924	0.0069	0.04117	0.0008	0.755
5	0.2876	0.0046	0.04086	0.0012	0.633

Table S-2 Mathematical equation for U-Pb data correction between standards and samples.

Rsam cor=Rsam mea*Rstd ref*{1/Rstd mea1* [1-(t^{sam}-tstd mea1)/(tstd mea2-tstd mea1)] +1/Rstd mea2*(t^{sam}-tstd mea1)/(tstd mea2-tstd mea1)}

Note: Rsam mea is the measured isotopic ratio of sample at time of t^{sam}, Rsam cor is the corrected isotopic ratio of the sample, Rstd ref is the reference isotopic ratio of the primary standard, and Rstd mea1 and Rstd mea2 are the measured isotopic ratios of primary standard at times of tstd mea1 and tstd mea2, respectively.

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U–Pb isotopic ratios are corrected with standard-sample bracketing method. Cassiterite AY-4 or other RMs was used as external standard, which was analyzed two times every 10 or 15 analyses of titanomagnetite samples, i.e. NIST612 +2 91500+2 AY-4 +2 PL-57 +2\sim3 YGX +10\sim15 samples+2 91500+2 AY-4 +2 PL-57 +2\sim3 YGX + NIST612 or NIST612 +2 91500+2 AY-4 +10\sim15 samples+2 AY-4 +2 91500 + NIST612. A short integrating time (\sim28 s) of external standard and samples were selected to correct the Pb/U fractionation and instrumental mass discrimination using ICPMSDataCal software. The U–Pb isotopic ratios of samples were calculated using a linear interpolation (with time) for every ten or fifteen analyses according to the variations of standard. Similar description of mathematical equation and the uncertainty calculation can be found in Tang et al. (2022, 2024)<sup>14,16</sup>.
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